Climate sensitivity and surface temperature

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Outline



- > Large uncertainty in GCMs \(\rightarrow \) simplified models
- * physical processes: clouds, precipitation, energy balance
- > Satellite & model studies:
 relationships among Ts, Rnet, external & internal forcings
- Forster & Gregory: JC 2006
 seasonal & annual means
 both external & internal forcings were considered.
- Lindzen and Choi: GRL 2009

 events, radiation changes
 ~1 years
- ❖ Spencer and Braswell: AGU 2009 & CERES STM 2009 certain Rnet & Ts phase states, chaotic system

few months

> Short-term variations of Ts & Rnet



Energy Balance Analysis



- > Energy balance:
- * TOA radiation changes
- surface temperature variations
- > Earth's heat reservoirs:
- * Ocean mixed layer
- Deep oceans
- Internal forcing -- heat transports among reservoirs
- > Sensitivity: $-\partial R/\partial T$

$$\mathbf{R}(\mathbf{F}, \mathbf{T}, \mathbf{N}) = \mathbf{R}_0 + (\partial \mathbf{R}/\partial \mathbf{F})\Delta \mathbf{F} + (\partial \mathbf{R}/\partial \mathbf{T})\Delta \mathbf{T} + (\partial \mathbf{R}/\partial \mathbf{N})\Delta \mathbf{N} + \mathbf{higher order terms}$$

 $(R_0 = 0; No N terms if entire climate system is considered)$

$$\mathbf{R} (\mathbf{F}, \mathbf{T}, \mathbf{N}) - (\partial \mathbf{R}/\partial \mathbf{F}) \Delta \mathbf{F} = (\partial \mathbf{R}/\partial \mathbf{T}) \Delta \mathbf{T} + (\partial \mathbf{R}/\partial \mathbf{N}) \Delta \mathbf{N}$$

$$\Delta \mathbf{R} = (\partial \mathbf{R}/\partial \mathbf{T})\Delta \mathbf{T} + (\partial \mathbf{R}/\partial \mathbf{N})\Delta \mathbf{N} \rightarrow \mathbf{empirical:} \Delta \mathbf{R} \text{ vs } \Delta \mathbf{T}$$



Background: Climate perturbation



$$Cp\frac{dTs}{dt} = (1 - \alpha)So - \varepsilon\sigma Ts^4$$

Cp: equivalent heat capacity

define:

 $\Delta Ts = T$

equilibrium state: $\Delta \alpha = \Delta \epsilon = 0$

$$Cp \frac{d\Delta Ts}{dt} = -\frac{4\varepsilon\sigma Ts^{4}}{Ts} \Delta Ts$$
$$= -\frac{4\times237}{288} \Delta Ts = -3.3\Delta Ts$$

fn = -3.3 Wm⁻²K⁻¹ (only for the equilibrium state)

At short time scales,
this feature is mixed with other processes.



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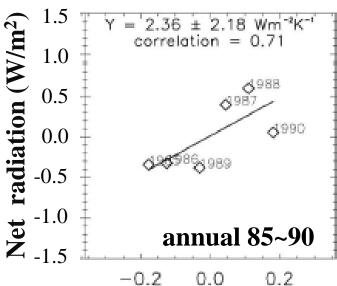
$$\mathbf{R} (\mathbf{F}, \mathbf{T}, \mathbf{N}) - (\partial \mathbf{R}/\partial \mathbf{F}) \Delta \mathbf{F} = (\partial \mathbf{R}/\partial \mathbf{T}) \Delta \mathbf{T} + (\partial \mathbf{R}/\partial \mathbf{N}) \Delta \mathbf{N}$$

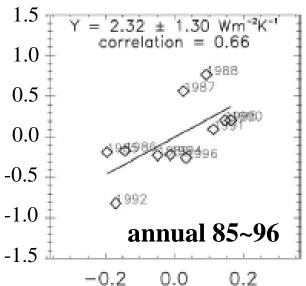
$$\Delta \mathbf{R} = (\partial \mathbf{R}/\partial \mathbf{T})\Delta \mathbf{T} + (\partial \mathbf{R}/\partial \mathbf{N})\Delta \mathbf{N} \rightarrow \mathbf{empirical:} \Delta \mathbf{R} \text{ vs } \Delta \mathbf{T}$$



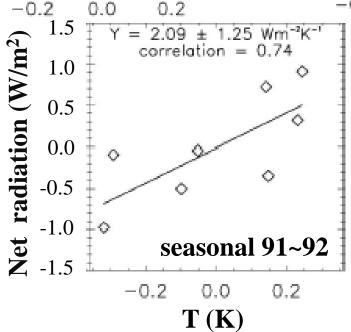
short-term relation







weak
positive
feedback



data for different time periods

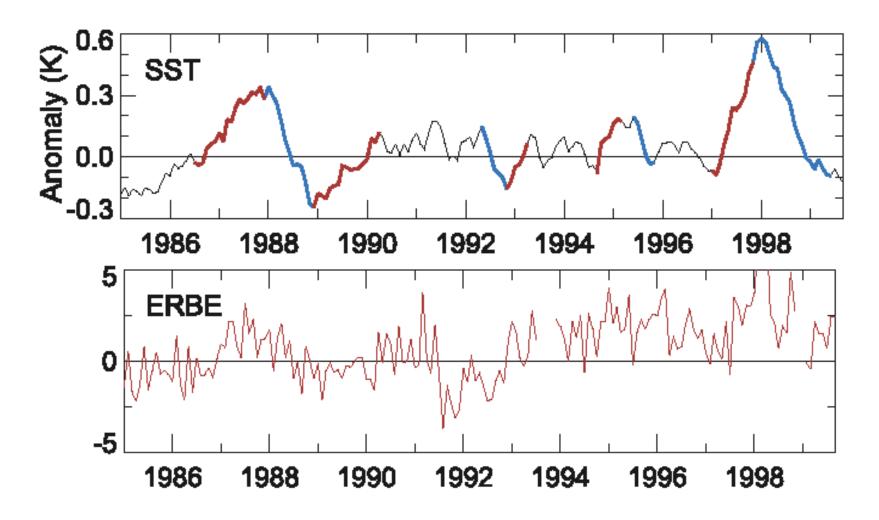
But ??: avg. N = 0; long-term feedback removed

Forster & Gregory: JC 2006



short-term relation: cont

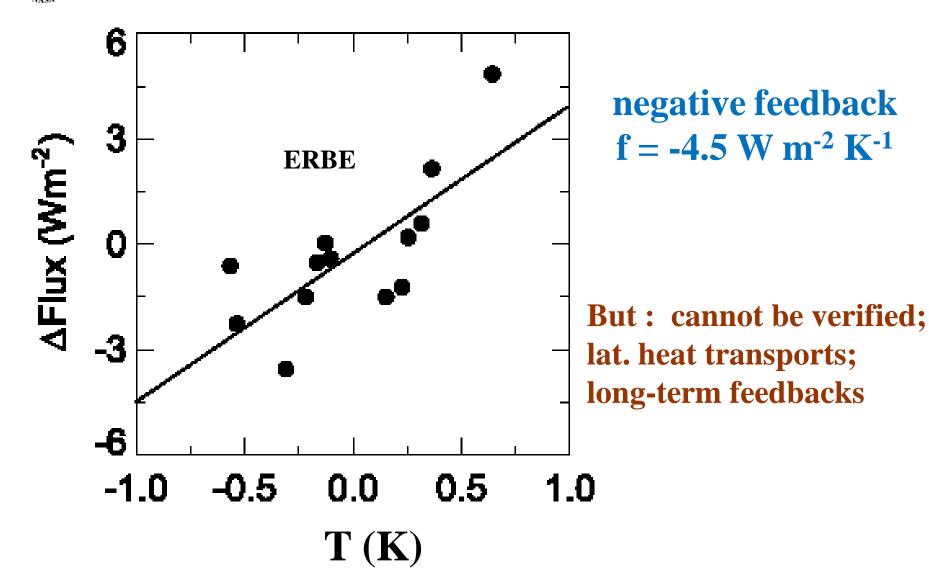






short-term relation: cont

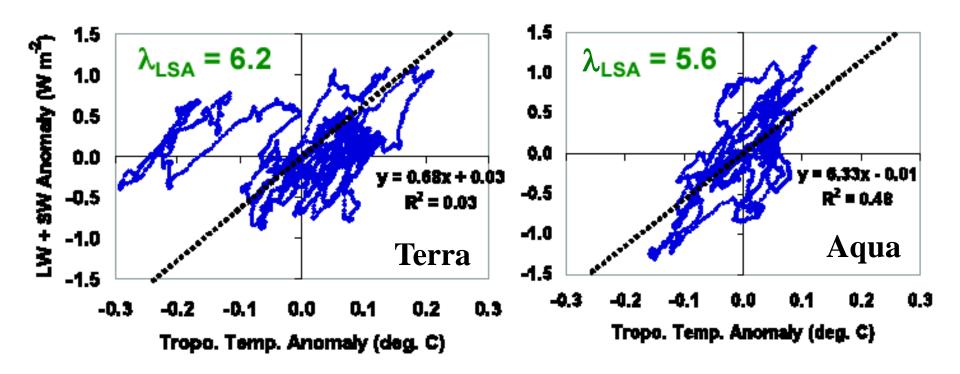






short-term relation: cont





Global Oceanic LW+SW Anomaly Total Feedback Parameter of ~6.0 W m⁻² K⁻¹



Observation Explanation



Perturbation model: energy balance

$$Cp\frac{dT}{dt} = F + ftotT + N + S$$

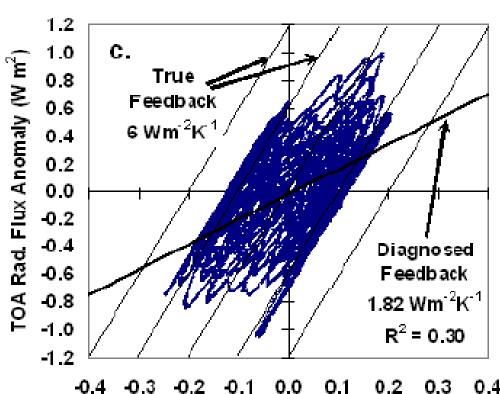
N: non-radiative heating (daily)

S: non-feedback natural radiative variability (5-yr cyc)

 f_{tot} : $f_n + f = -6 \text{ Wm}^{-2}\text{K}^{-1}$

F: F = 0 or removed

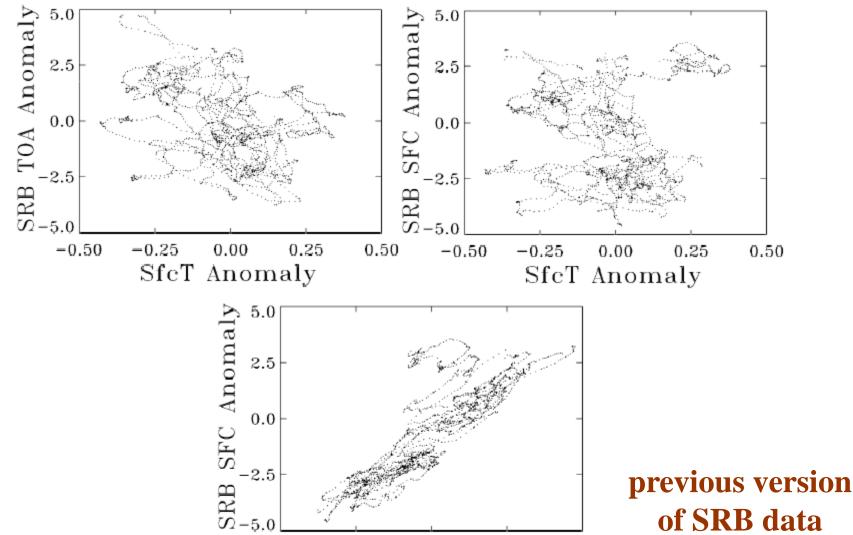
Cp: 50 m water





SRB Results





0.0

SRB TOA Anomaly

-5.0

-2.5

2.5

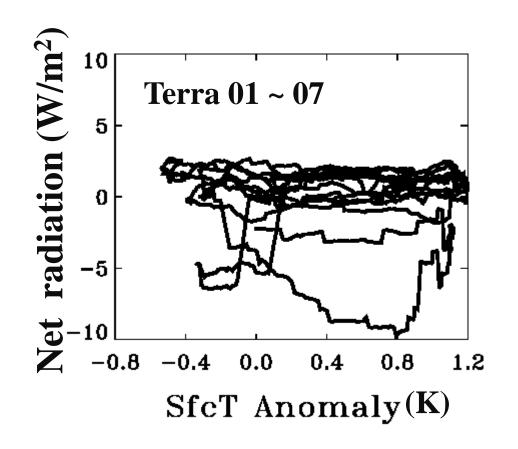
5.0

of SRB data



CERES observations





90-day running mean

some indication of short-term responses of radiation to sfc T

Sfc Temp from CERES MOA



Analysis Approach

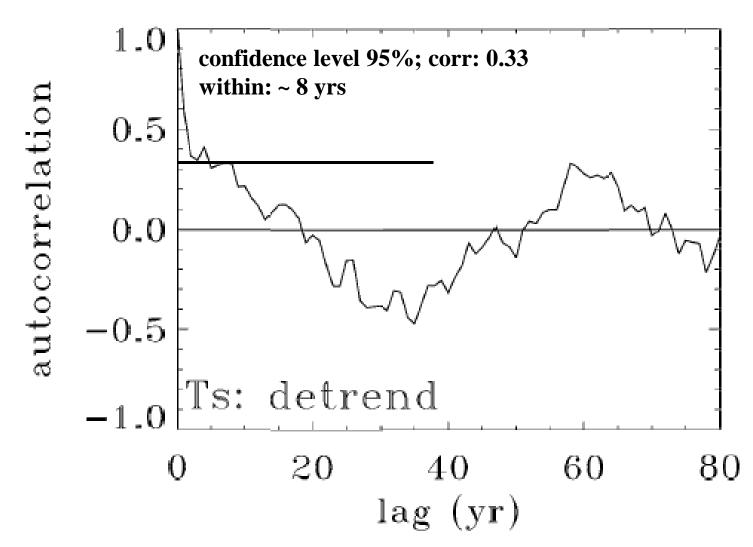


- > Energy balance:
- * Spencer and Braswell: 2009 (AGU & CERES STM)
- ❖ Lin et al. 2010 ACP
- > Basic physical components:
- * ocean mixed layer
- internal and external forcing
- * TOA radiation
- > Additional considerations:
- climate system memory
- different feedback tests



System memory





Ts: detrended GISS sfc T



Modeling Considerations



Perturbation theory: energy balance model

$$Cp\frac{dT}{dt} = F + fsT + N + S + \frac{fm}{t_0} \int_{t-t_0}^{t} Tdt'$$

N: non-radiative heating (daily) \leftrightarrow avg N = 0

S: non-feedback natural radiation (5-yr cycle) \leftrightarrow avg S = 0

$$f_s$$
: $f_s = f_n + f = -6 \text{ Wm}^{-2}\text{K}^{-1}$; $f = -2.7 \text{ Wm}^{-2}\text{K}^{-1}$

$$f_{tot}$$
: $f_{tot} = f_s + f_m$

F: F = 0 or removed

to: memory length \leftrightarrow minimal (1 year); other lengths also tested

Cp: 100 m mixed layer ocean (slab ocean)

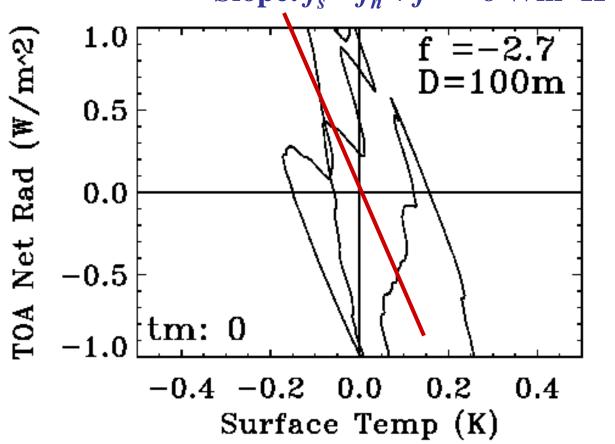
last 10-year results of 100-year run



System without memory







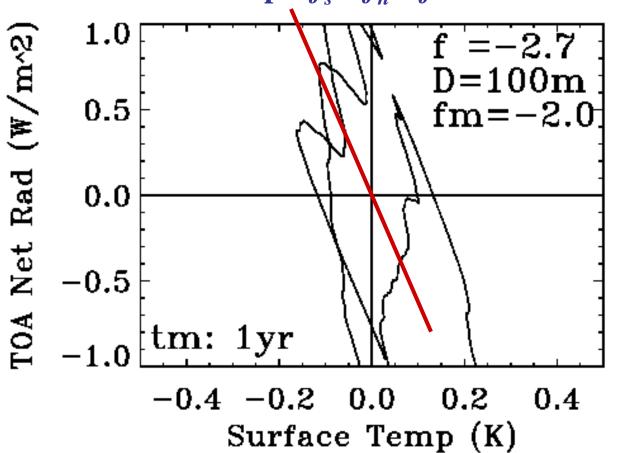
Similar results as previous studies



memory system



Slope: $f_s = f_n + f = -6 \text{ Wm}^{-2}\text{K}^{-1}$

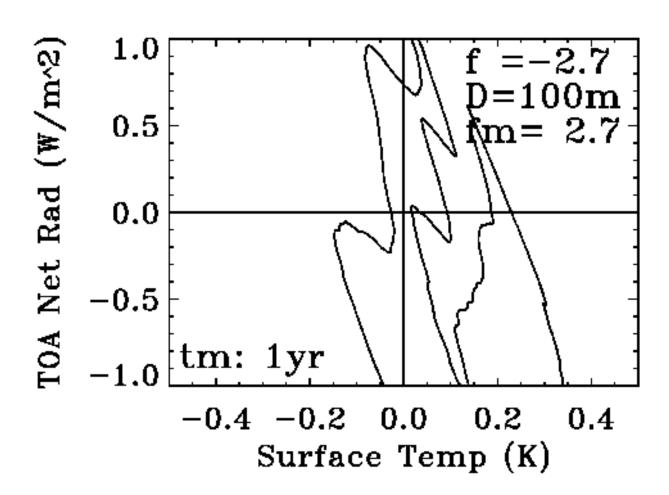


Extreme strong negative feedback system Total feedback parameter: $f_{tot} = f_s + f_m = -8 \text{ Wm}^{-2}\text{K}^{-1}$



memory system





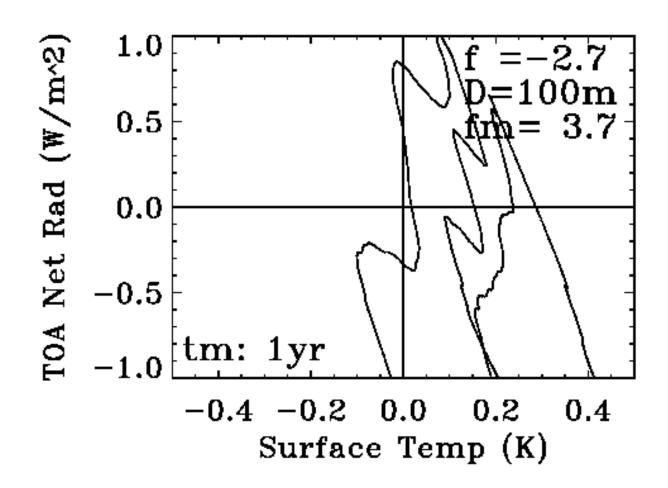
Neutral feedback system

Total feedback parameter: $f_{tot} = f_s + f_m = -3.3 \text{ Wm}^{-2}\text{K}^{-1}$



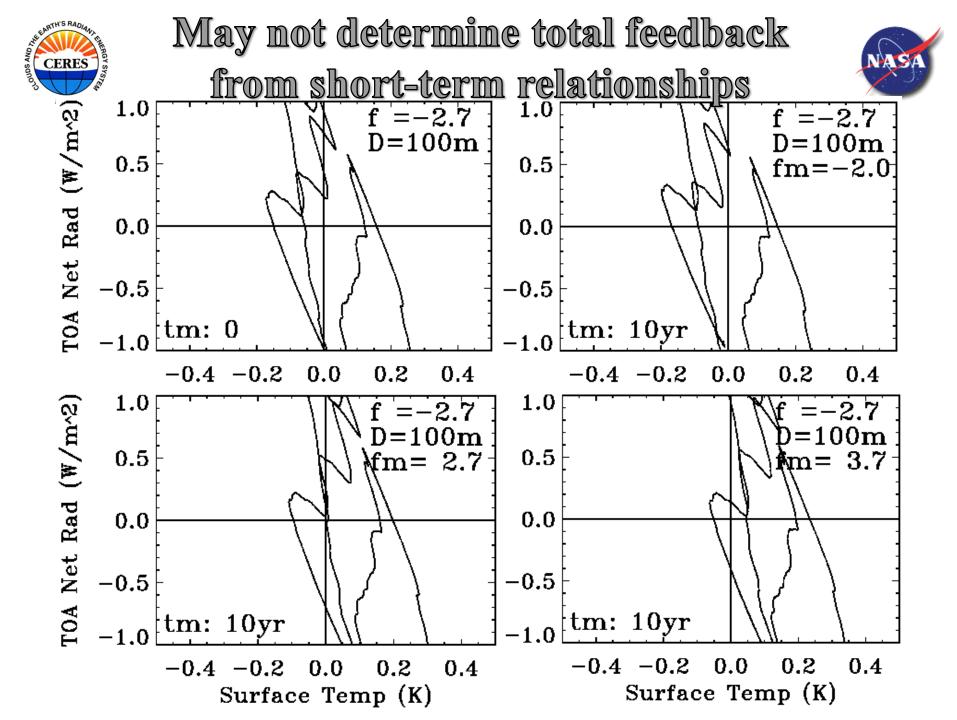
memory system





Slightly positive feedback system

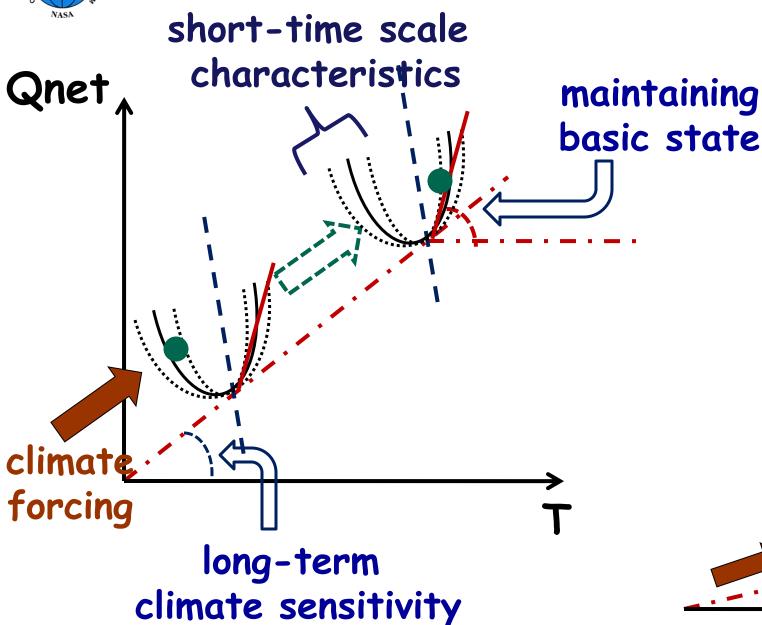
Total feedback parameter: $f_{tot} = f_s + f_m = -2.3 \text{ Wm}^{-2}\text{K}^{-1}$





climate sensitivity







Summary



- Energy balance model for explanation of the anomolies of TOA net radiation and surface temperature.
- Major physical processes of the climate system, such as internal and external forcing, and system memory, are considered.
- Internal non-radiative heating is needed due to slab ocean approximation (no vertical heat transport) and chaotic feature of the climate.
- Cannot use short-time relationships between sfc temp and net radiation of the climate system to mimic the feedbacks of climate change: different scales, different physics



Acknowledgement



Many people, especially David Young, Gary Gibson, and Don Garber, have significant supports for this study.